

SALINITY AND WATER TEMPERATURE ASSESSMENT OF THE TIDAL MARSHES FROM THE W PORTUGUESE COAST, AS AN ECOLOGICAL TOOL TO PALAEOENVIRONMENTAL RECONSTRUCTIONS BASED ON FORAMINIFERA AND OSTRACODA ASSEMBLAGES

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Abstract:

A seasonal study of temperature and salinity of estuarine and sediment interstitial water of tidal marshes was undertaken along three estuaries of W Portuguese coast (Minho, Tejo and Mira). The climatic N-S transition from wet Atlantic to Mediterranean features appear clearly imprinted in the distribution of tidal marsh assemblages, like foraminifera and ostracoda, mainly reflecting the water salinity gradient control. The Minho low estuary tidal marsh tends to be flooded by estuarine water ranging from 0.5‰ to 32‰ in each tide cycle, even during dry seasons. However, the marsh hydrological balance sustains a more stable environment where the salinity of interstitial water measurements yielded 8‰ to 16‰. In contrast the Tejo and Mira salt marsh flooding waters record a narrow range between 33‰ and 36‰, in spring, and between 29‰ and 36‰ in autumn. The climatic control of evaporation/precipitation balance produces an enhanced salinity of marsh interstitial water, that can reach hypersaline conditions, with maximum records of 53‰ in Tejo and 48‰ in Mira lower estuaries. These environmental differences along the W Portuguese coast are recorded by the tidal marsh assemblages, namely foraminifera and ostracoda. In the low salinity Caminha salt marsh, living foraminifera are essentially composed by the agglutinated species *Haplophragmoides manilaensis*, *Miliammina fusca*, *Pseudothurammina limnetis*, *Psammosphaera* sp. and *Trochammina salsa*. The modern ostracoda assemblage includes *Leptochytère baltica*, *Leptochytère psammophila*, *Leptocythere* sp. A and *Tuberoloxoconcha* sp.1. In the Tejo and Mira salt marsh *Ammonia beccarii*, *Ammonia tepida*, *Haynesina germanica*, *Jadammina macrescens*, *Trochammina inflata*, are the dominant foraminifera and *Loxoconcha malcomsoni*, *Terrestricythere* cf. *elisabethae*, *Tuberoloxoconcha* cf. *atlantica* and *Xestoleberis labiata* prevail as well as many other more marine ostracoda species, such as *Basslerites teres* and *Leptocythere fabaeformis*. This study highlights that the knowledge of driven ecological parameters of modern assemblages (usually preserved in fossil record), is fundamental to support reliable paleoclimatic and palaeoenvironmental reconstructions.

Keywords: salinity, tidal marshes, foraminifera, ostracoda, Portuguese W coast

Resumo:

Avaliação da salinidade e da temperatura da água nos sapais da costa Oeste de Portugal, na perspectiva da reconstrução paleoambiental com base na ecologia de Foraminíferos e Ostracodos

O registo sazonal da temperatura e da salinidade da água dos estuários e da água intersticial dos sedimentos de sapal nos rios Minho, Tejo e Mira, durante a enchente, integram o estudo das associações de foraminíferos e ostracodos actuais da costa Oeste de Portugal. Nestes trabalhos podemos observar a transição entre o padrão climático húmido do NO da Península Ibérica e o padrão mediterrâneo a SO.

O sapal do baixo estuário do rio Minho tende a ser inundado por águas com grande amplitude de salinidade, de 0.5‰ a 32‰ em cada ciclo de maré, que persiste mesmo nas estações mais secas. No entanto o balanço hidrológico do sapal contribui para um ambiente relativamente estável, onde os valores da salinidade da água intersticial variam entre 8‰ e 16‰.

No baixo estuário do rio Mira, a salinidade das águas que inundam os sapais durante a preia mar varia apenas entre 33‰ e 36‰, na Primavera, e entre 29‰ e 36‰ no Outono. O balanço da evaporação/precipitação no contexto climático do Sul, produz um aumento da salinidade da água intersticial até à hipersalinidade, que atinge 53‰ nos sapais do Tejo e 48‰ nos do Mira.

Estas diferenças das condições ambientais ao longo da costa Oeste de Portugal estão reflectidas na composição das associações dos organismos de sapal, nomeadamente dos foraminíferos e dos ostracodos. Sob as condições de baixa salinidade registadas no sapal de Caminha, a biocenose de foraminíferos é composta essencialmente pelas espécies de carapaça aglutinada *Haplophragmoides manilaensis*, *Miliammina fusca*, *Pseudothurammina limnetis*, *Psammosphaera* sp. e *Trochammina salsa*. A associação dos ostracodos actuais inclui as espécies *Leptochytère baltica*, *Leptochytère psammophila*, *Leptocythere* sp. A e *Tuberoloxoconcha* sp. 1. Nos sapais do Tejo e do Mira as espécies *Ammonia beccarii*, *A. tepida*, *Haynesina germanica*, *Jadammina macrescens*, *Trochammina inflata*, dominam a associação de foraminíferos, enquanto *Loxoconcha malcomsoni*, *Terrestricythere* cf. *elisabethae*, *Tuberoloxoconcha* cf. *atlantica* e *Xestoleberis labiata* se destacam numa associação de ostracodos com diversidade de espécies de características mais marinhas, como *Basslerites teres* e *Leptocythere fabaeformis*.

A relação encontrada entre a salinidade e a composição das associações de foraminíferos e ostracodos, mostra que a caracterização dos factores ecológicos e das biocenoses (dos grupos que integram o registo fóssil) em análogos actuais, são determinantes para a reconstituição fiável da evolução paleoclimática e paleoambiental.

Palavras-Chave: salinidade, sapais, foraminíferos, ostracodos, Costa Oeste de Portugal

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1. INTRODUCTION

Estuaries are highly productive environments but they also undergo intense impacts of human activities during the last century. The intertidal fringe is often occupied by salt marshes that represent one of the most suitable environments to search for the record of regional and global forcing mechanisms. Actually the almost constant sedimentation, that builds these zones, often represents the most continuous post-glacial geological record available on shore (e.g. DAVIS & FITZGERALD 2004).

Since the 1950s several researchers studied salt marshes aiming to define environmental proxies, like foraminifera and ostracoda, to be used in palaeoclimatic and palaeogeographic interpretations, namely within the context of global change evaluation (e.g. PHLEGER & WALTON 1950; PUJOS 1971; SCOTT & MEDIOLI 1980; GEHRELS 1994; DE RIJK 1995; HAYWARD *et al.* 1999; CEARRERA *et al.* 2000, 2007; HORNE & BOOMER 2000; SEN GUPTA 2002; HORTON & EDWARDS 2006; MURRAY 2006; CABRAL & LOUREIRO, 2013). Such applications of microfossils record rests upon the Principle of Actualism which states that the composition of the micro-organisms assemblages strongly depends on the environmental biotic and abiotic parameters thereby making them reliable proxies. Consequently, their interpretation must be supported by a robust and site specific data base built upon the study of temperature, salinity, pH, dissolved oxygen, and CaCO₃ content (among others) that constrains the distribution of modern living communities (e.g. MORENO *et al.* 2005, 2007; LEORRI *et al.* 2008; FATELA *et al.* 2009; LOUREIRO *et al.* 2009; VALENTE *et al.* 2009; CABRAL & LOUREIRO 2013).

In this paper we present the results of synoptic measurements of estuarine and interstitial waters temperature and salinity seen as main ecological parameters to the control of living tidal marsh assemblages along the West Portuguese coast (e.g. MORENO *et al.* 2006, 2007; FATELA *et al.* 2007, 2009; CABRAL & LOUREIRO 2013).

2. REGIONAL SETTING

The W coast of Portugal faces the NE Atlantic and develops approximately from 37° 00'N to 41° 50'N and between 8° 40'W and 9° 30'W, over more than 700km long, often interrupted by many estuaries. The estuaries of Minho, Tejo (Tagus) and Mira were selected from those where preserved tidal marshes can be found in the lower estuary section.

Tides present a semi-diurnal high-mesotidal regime along this coast. The tidal range varies between 2m, during neap waters, and almost 4m in spring waters but the astronomical tide levels are often incremented by storm surges (TABORDA & DIAS 1991).

The Minho River defines the political border with Spain along 77km, joining the regions of Minho and Galicia (Fig. 1) and discharging the

sea through a barred estuary trending NNE-SSW. The dynamic tide is felt up to 40km upstream, due to the large tidal range and to both the smoothness and low gradient of the Minho's outlet (ALVES 1996). The upstream limit of marine water influence is not consensual and disparate distances of 9 to 35km have been found (BETTENCOURT *et al.* 2003; MORENO *et al.* 2005).

The Minho watershed spans over 17 080km² of igneous and metamorphic rocks basement, draining the rainiest region of Portugal. The average annual precipitation is ca. 1 600mm but exceptionally may reach 3 500mm (FATELA *et al.* 2014). The yearly average fluvial discharge is about 300m³/s and the winter peak discharge (December to March) usually exceeds 1 000m³/s (BETTENCOURT *et al.* 2003).

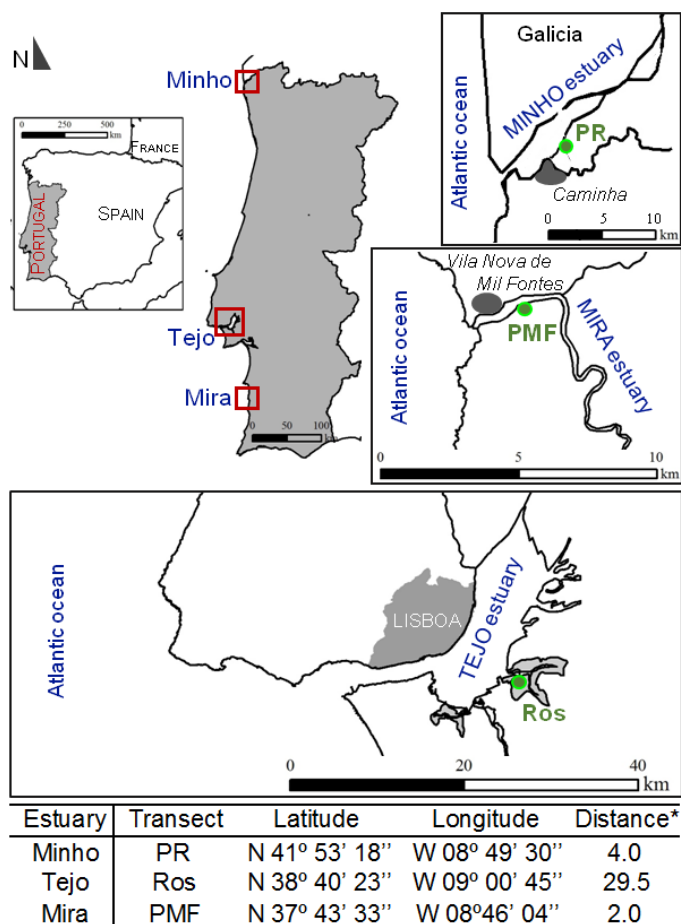
The Tejo river has one of the largest estuaries of Western Europe, covering more than 325km². Forty percent of this area emerges during low water spring. The hydrographical basin develops in Portugal and Spain over 81 310km², occupied by a large variety of igneous, metamorphic, calcareous and detritic rocks. The average annual precipitation in the Portuguese sector of Tejo basin reaches ca. 900mm. The fluvial average discharge is around 300m³/s, but may range out between 250m³/s and 5400m³/s under extreme dry or wet conditions (BETTENCOURT *et al.* 2003; ARH_Tejo 2011).

The dynamic tidal effects are felt up to a distance of 80km upstream and the salt edges reaches 50km upstream under the average flux of 300m³/s (BETTENCOURT *et al.* 2003).

Mira is an important river of SW Portugal whose hydrographical basin occupies an area of 1600km² of mostly greywakes, pelites, slates, schists and conglomerates. The estuary extends from the mouth to SSE for almost 40km of incised meanders, passing Odemira, where the dynamic effect of tide completely felts down (ICNB 2008). Annual precipitation is around 645mm, with a monthly average between 2.6mm and 103mm recorded in a series of 75 years. However, 0mm is frequently record in summer (ICNB 2008). The fluvial discharge ranges from 0m³/s during dry summers to 500m³/s in winter and spring rainy periods, leading to a yearly average fluvial discharge of 2.9m³/s. Tides are thus the main flow component in the Mira estuary (INAG 2011).

3. METHODS

Salinity and temperature of estuarine and sediment pore-water were measured along 3 low estuary transects in Minho (Pedras Ruivas - PR), Tejo (Rosário - Ros) and Mira (Ponte de Vila Nova de Mil Fontes - PMF) tidal marshes (Fig. 1), under spring and autumn season conditions. The interstitial water was allowed to seep and accumulate inside perforated PVC tubes previously inserted into the sediment to a depth of 40cm below the surface, following De Rijk (1995), and covered with aluminium foil



* Distance to river mouth (km)

Fig. 1. Location of studied salt marsh transects along the w coast of Portugal, in the estuaries of Minho (PR – Pedras Ruivas), Tejo (Ros – Rosário) and Mira (PMF – Ponte de Vila Nova de Mil Fontes).

Fig. 1. Localização dos sapais estudados ao longo da costa oeste de Portugal, nos estuários dos rios Minho (PR – Pedras Ruivas), Tejo (Ros – Rosário) and Mira (PMF – Ponte de Vila Nova de Mil Fontes).

between measurements. The salinity and temperature were also measured every 15 minutes, close to the marsh surface during the rising tide at PR (Minho) and PMF (Mira) transects. These measurements started just after the beginning of surface submersion and ended at high water slack, unless interrupted for safety reasons. Simultaneous measurement of the salinity and temperature of estuarine water has also been made close to the channel bottom, in high and low water during spring tide. Water parameters were measured using a multiparameters Horiba U-22, a WTW LF 191 and 197i probes. The water salinity terms of limnetic, oligohaline, mesohaline, polyhaline, euhaline and hypersaline, used hereafter, follow the criteria defined by the Venice System (1959).

The altimetric data of the sampling profiles have been obtained using a Zeiss Elta R55 total station from a bench mark connected to the national altimetric datum (Cascais), using the GPS differential positioning combined with a regional geoid model and linked to local chart datum (e.g. FATELA *et al.* 2009).

4. RESULTS

4.1. Marsh sediment interstitial water

4.1.1. Minho estuary

The temperature of Caminha marsh interstitial waters (PR) tend to have a stable pattern from tidal flat to high marsh. The main differences were found between spring (average: 11.5°C) and autumn (average: 16.4°C) temperatures, as shown in Table 1.

The salinity measurements in the PR transect (Fig. 2), under spring season conditions, yielded 0.9‰ in the subtidal domain, contrasting with low marsh values of 15.8‰ (PR5) and 7.7‰ of high marsh (PR10). This trend is maintained in autumn, but an increase in salinity values becomes clear (Fig. 2) with 1.9‰ in the subtidal domain and 9.4‰ in the high marsh (PR10). The maximum salinity of PR5 is the same as in spring but the other records of low marsh also arise (Tab. 1; Fig. 2).

Table 1. Temperature and salinity of marsh sediment interstitial water and estuarine water. PR – Pedras Ruivas; Ros - Rosário; PMF - Ponte de V.N. Mil Fontes; LT - low tide; HT - high tide.
Tabela 1. Temperatura e salinidade da água intersticial do sedimento de sapal e da água de estuário. PR – Pedras Ruivas; Ros - Rosário; PMF - Ponte de V.N. Mil Fontes; LT – Baixa Mar; HT – Preia Mar.

Temperature °C													
		PR1	PR2	PR3	PR4	PR5	PR6	PR7	PR8	PR9	PR10	LT	HT
MINHO	Spring	11.9				11.7	12.1	11.3	10.3				
	Autumn					16.8	17.3	16.0	15.7	15.9			
TEJO	Spring	Ros1	Ros2	Ros3	Ros4	Ros5	Ros6	Ros7	Ros8	Ros9	Ros10	LT	HT
	Autumn	19.7	15.2	14.4	14.0	14.5	12.6	14.1	14.3	16.1	13.7	13.6	13.0
MIRA	Spring	PMF1	PMF2	PMF3	PMF4	PMF5	PMF6	PMF7	PMF8	PMF9	PMF10	LT	HT
	Autumn	16.3	16.0	15.2	15.1	14.9	14.5	14.7					18.2
Salinity ‰													
		PR1	PR2	PR3	PR4	PR5	PR6	PR7	PR8	PR9	PR10	LT	HT
MINHO	Spring	0.9				15.8	11.9	8.4	7.7				
	Autumn					15.8	14.4	12.2	10.4	9.4			
TEJO	Spring	Ros1	Ros2	Ros3	Ros4	Ros5	Ros6	Ros7	Ros8	Ros9	Ros10	LT	HT
	Autumn	28.0	31.6	33.2	33.1	39.5	31.8	28.4	52.9	45.7	23.6	29	
MIRA	Spring	PMF1	PMF2	PMF3	PMF4	PMF5	PMF6	PMF7	PMF8	PMF9	PMF10	LT	HT
	Autumn	33.8	33.7	35.9	42.0	47.5	28.8	11.7					32.8

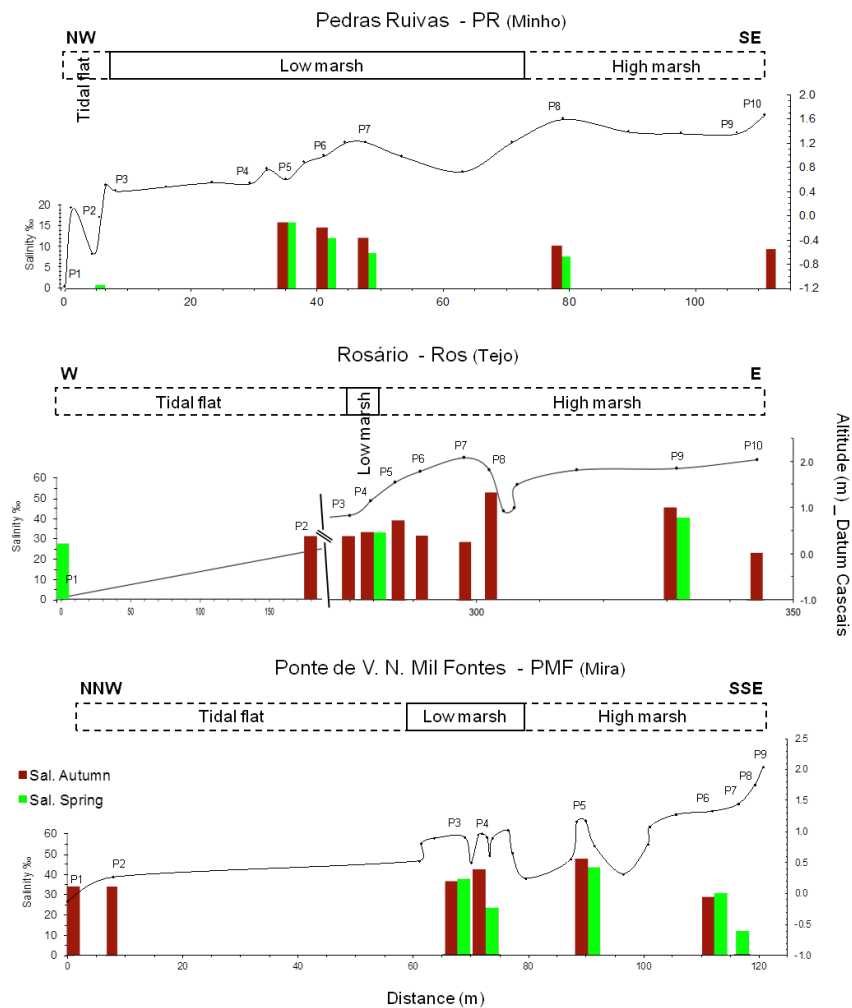


Fig. 2. Sampling profiles of studied transects of Pedras Ruivas (PR – Minho), Rosário (Ros – Tejo) and Ponte de V.N. Mil Fontes (PMF – Mira), showing the measured spring and autumn salinity of sediment interstitial water.
Fig. 2. Valores da salinidade da água intersticial do sedimento de sapal medidos, na Primavera e no Outono, nos perfis de amostragem de Pedras Ruivas (PR – Minho), Rosário (Ros – Tejo) e Ponte de V.N. Mil Fontes (PMF – Mira).

4.1.2. Tejo estuary

At Rosário salt marsh the temperature of interstitial waters measured across the Ros transect is also very stable. The main differences are also related to season but reveal a reverse trend, when compared with Minho tidal marsh. The average temperature is higher in spring (18.1°C) than in autumn (14.0°C; Table 1).

The salinity measurements in the Ros transect (Fig. 2), under spring season conditions, yielded 28‰ in the tidal flat (Ros1), 32.9‰ in the low marsh (Ros4) and 40.4‰ in the high marsh (Ros9). This direct relation of increasing salinity with altitude of the marsh surface is also clear in autumn (Fig. 2), ranging from 31.6‰ in the tidal flat (Ros2) and 52.9‰ in the high marsh (Ros8). However, a significant drop to 23.6‰ is recorded at Ros10, in highest high marsh zone (Tab. 1).

4.1.3. Mira estuary

The stability of temperature in the marsh interstitial waters is also observed along the PMF transect (Tab. 1), with no significant difference between spring (average: 16.2°C) and autumn (average: 15.3°C) records.

The spring season salinity measurements in the PMF transect (Fig. 2), yielded 32.8 ‰ in the subtidal domain, an average of 30.1‰ in the low

marsh and a maximum of 43.3‰ (PMF5) in the high marsh. The relation of increasing salinity with altitude of the marsh surface is also clear in this transect, rising in autumn (Fig. 2), except in the subtidal domain (33.8‰). Average salinity record rises to 38.9‰ in the low marsh and reaches new maximum of 47.5‰ (PMF5) in the high marsh. In both seasons a significant drop to 11.7‰ and 28.8‰, respectively, is recorded at the transition to highest high marsh zone (Tab. 1). All the sampling points from the highest high marsh were dry, in both spring and autumn campaigns, avoiding the measurement of interstitial water parameters.

4.2. Tidal marsh flooding water

4.2.1. Minho estuary

The Minho lower estuary records a wide range salinity twice a day (MORENO *et al.* 2005) resulting from the balance between river flow and tidal regime. The measurements close to the sediment surface across the PR transect, during its submersion by a spring tide, show that from the arrival of estuarine water at each sampling point until the complete flooding of the marsh, the salinity values at the sediment – water interface ranged from 0.5‰ to 31.6‰ in the tidal flat, 3.1‰ to 31.6‰ in the low marsh and 7.0‰ to 23.1‰ in the high marsh (Fig. 3).

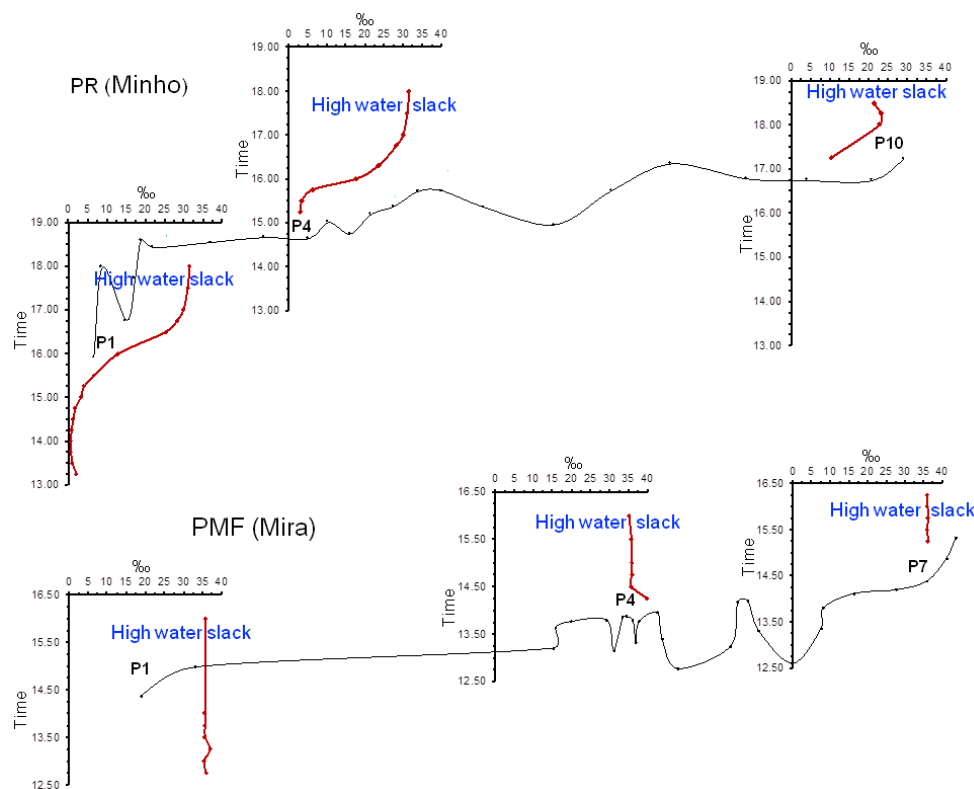


Fig. 3. Example of water salinity variation along tidal flat, low marsh and high marsh, during the spring tide flooding, in the Minho and Mira low estuaries.

Fig. 3. Exemplo da variação da salinidade ao longo da superfície do raso de maré, do baixo sapal e do alto sapal, durante a enchente duma maré de águas vivas, no estuário inferior dos rios Minho e Mira.

4.2.2. Mira estuary

In the Mira lower estuary the tidal regime do not force a marked wide range of salinity. Our records show that in autumn the salinity ranges between 28.6‰ and 35.5‰, and between 32.8‰ and 35.8‰ in spring season.

The measurements performed during a rising spring tide across the PMF transect, show that from the arrival of estuarine water at each sampling point until the complete flooding of the marsh, the salinity values at the sediment – water interface range from 35.4‰ to 36.0‰ in the tidal flat, 35.0‰ to 39.8‰ in the low marsh and 35.4‰ to 36.5‰ in the high marsh (Fig. 3).

5. DISCUSSION

A distinctive attribute of the W coast of Portugal is the transition from the warm and rainy Atlantic to the hot and dry Mediterranean climate across a distance of around 400km, in a narrow range of latitude. That means that estuarine environments and hydrographic basin are constrained by a Cfb climate type (temperate, with warm summer and no dry season) at North, and Csa climate type (temperate with hot and dry summers) at South (PEEL *et al.* 2007).

The temperature of marsh sediment interstitial waters is quite stable across all the transects for the same site and campaign. So when differences appear, they are directly connected with seasonality of climate in the temperate regions. The climatic fingerprint is also noted in the North to South increasing temperatures when the Caminha records are compared with Tejo and Mira marsh interstitial waters.

In the northern estuary of Minho River, the penetration of salt edge faces the resistance of lower basin morphology and of hydrological features resulting from the climate type mentioned above. On one hand, a widespread siltation has led to a very shallow estuary, with a very limited volume to hold the tidal prism. On the other hand, the intense and long precipitation regime generates a persistent run-off. In Summer, under spring tide conditions, the euhaline water (30‰ to 35‰) penetrates less than 5km upstream during high tide (MORENO *et al.* 2005). However, at low tide the sea water is completely flush out, with prevailing oligohaline conditions (0.5‰ to 5‰) all over the lower estuary (Fig. 4). In the rainy spring time, at spring high tide, the same domain is occupied by 29‰ to 26‰ waters (polyhaline), from the mouth to less than 5km upstream. During low tide freshwater prevails and the most part of the estuary acquires limnetic conditions with 0‰ to 0.5‰ (Fig. 4). Nevertheless, the daily salinity dynamic stress is considerably smoothed inside the marsh sediment. The salinity of tidal flooding waters is retained in the sediment pore water according to

the submersion time, freshwater seepage from upland and evaporation/precipitation balance. In the Caminha tidal marsh the highest salinity (mesohaline) is found in the transition from low to high marsh zone (Fig. 3) as a result of a longer submersion with euhaline water. At the high water slack the top of high marsh tends to be flooded by polyhaline water, resulting from the less dense brackish mixture from Coura drainage, that is pushed onto the upland margin of Caminha tidal marsh. Moreover, interstitial high marsh water salinity is more affected by dilution with precipitation and freshwater inflow from surrounding land and groundwater.

In the Tejo and Mira marsh sediment interstitial waters the salinity values are higher than in Caminha tidal marsh. The low marshes of those southern estuaries are impregnated with euhaline waters that became hypersaline in the high marsh. There is a clear direct relation between pore-water salinity and marsh surface altitude. However, altitude is not a factor by itself but reflects the true process reducing the time submersion by tidal flooding waters from tidal flat to the high marsh. After the tidal retreat the higher zones are exposed for a longer period where evaporation prevails due to the hotter and dryer season's conditions felt in southern estuaries.

The Mira lower estuary tends to preserve a body of euhaline water even in low tides (BLANTON & ANDRADE 2001). Under spring season conditions, low marsh may be covered by hypersaline water in the beginning of the flooding tide, tending to the euhaline conditions in the high tide (Tab. 1; Fig. 3). This record reflects the importance of evaporation during low tide and the consequent precipitation of salt on the tidal flat and marsh surfaces. The rising tide dissolves this precipitated salt increasing significantly the salinity of the early flooding waters. By contrast in the Caminha tidal marsh we have recorded a gradual salinity increase, from limnetic to euhaline, in tidal flooding waters.

The Tejo and Mira salt marshes also record a relative decrease of pore-water salinity in the highest high marsh, suggesting that dilution by freshwater inflow from surrounding land and groundwater in the upland transition may be present under the dryer conditions of Csa climate type.

A seasonal cycle was found in the salinity of marsh interstitial water that leads to salt concentration in sediment during the dryer and hotter months (autumn records), and a dilution during the rainy and colder months (spring records), reflecting the evaporation/precipitation balance along the W coast.

Salinity is very important for the metabolism of organisms, namely by osmosis (MURRAY 2006) and for the production of carbonated tests. This influence is very clear in the composition and distribution of foraminifera and ostracoda assemblages. They may tolerate a wide range of salinity,

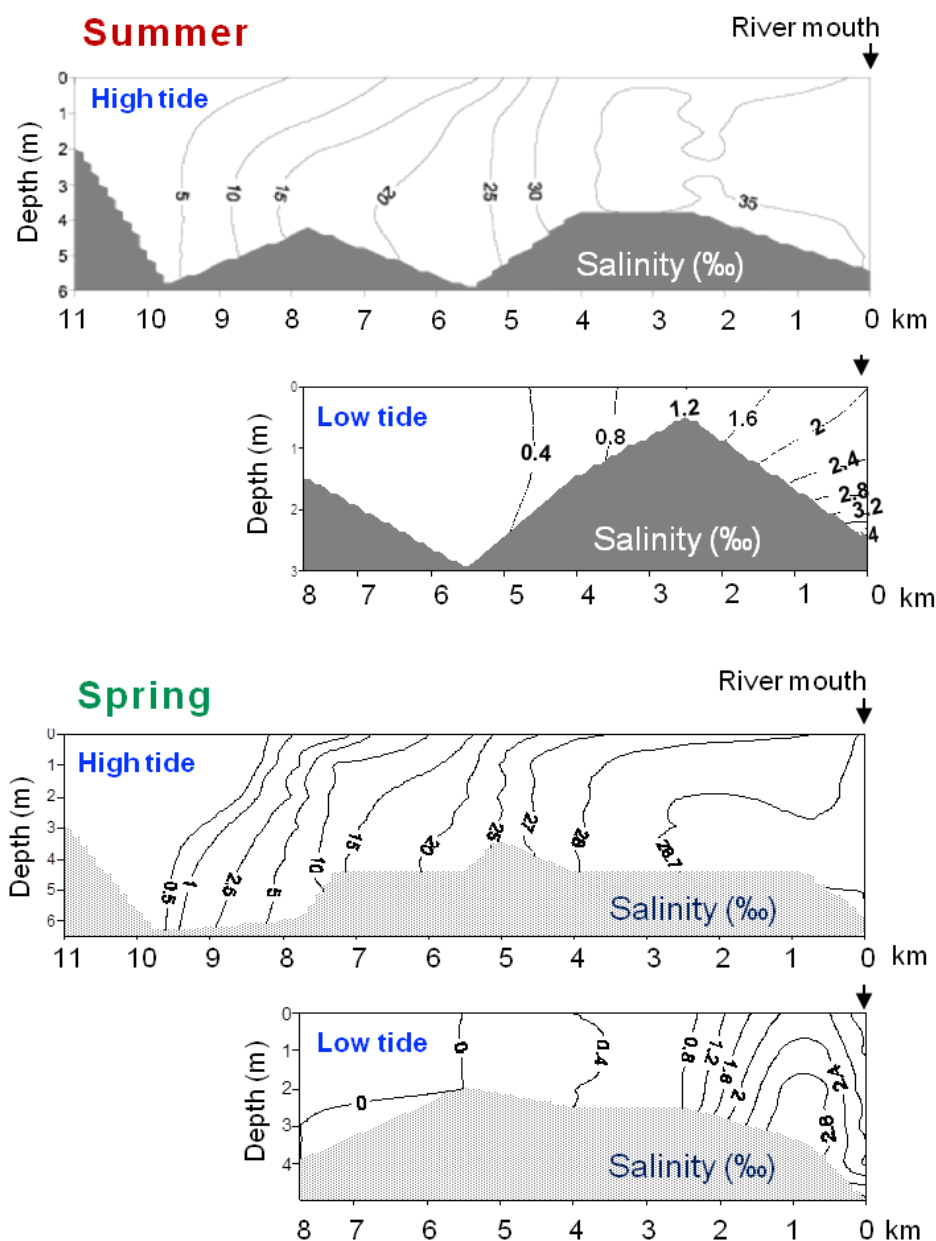


Fig. 4. Seasonal salinity profiles of Minho low estuary under summer and spring season conditions.
 Fig. 4. Perfis sazonais da salinidade medida no estuário inferior do rio Minho, em situação de Verão e de Primavera.

from limnetic to hypersaline conditions, but most species are found under euhaline conditions where the assemblages exhibit the highest diversity (e.g. HAQ & BOERSMA 1978; BOOMER & EISENHAUER 2002; SEN GUPTA 2002; FRENZEL & BOOMER 2005; MURRAY 2006).

The salt marsh foraminiferal assemblages under low salinity conditions are essentially composed by agglutinated species. For instance the Caminha salt marsh assemblages of living foraminifera include species like *Haplophragmoides manilaensis* Andersen, 1953, *Miliammina fusca* (Brady, 1870), *Pseudothurammmina limnetis* (Scott and Medioli, 1980), *Psammospaera* sp., *Trocham-*

minita salsa/ irregularis (Cushman and Brönnimann, 1948) and *Polysaccamina ipohalina* Scott, 1976. When the salinity of marsh sediment rises till polyhaline to hypersaline conditions, as found in Tejo and Mira estuaries, they are replaced by *Ammonia beccarii* (Lineaus, 1758), *A. tepida* (Cushman, 1926), *Haynesina germanica* (Ehrenberg, 1840), *Jadammina macrescens* (Brady, 1870), *Trochammina inflata* (Montagu, 1808), *Tiphotrocha comprimata* (Cushman and Brönnimann, 1948), *Paratrochammina guaratibaensis* (Brönnimann, 1986) and *Siphotrochammina lobata* Saunders, 1957, among other minor species (e.g. HAYWARD *et al.*, 1999; DEBENAY *et al.*, 2000; SCOTT *et al.*, 2001; SEN

GUPTA, 2002; MORENO *et al.*, 2006; FATELA *et al.*, 2009, 2014). The same trend is clearly shown by Cabral and Loureiro (2013), among other authors, in the distribution of modern ostracoda assemblages. For instance, the brackish species found in Caminha tidal marsh, like *Leptocythere baltica* Klie, 1929, *Leptocythere psammophila* Guillaume, 1976, *Leptocythere* sp. A and *Tuberoloxoconcha* sp. 1 are replaced by a diverse assemblage in the Tejo and Mira salt marshes, under euhaline to hypersaline conditions, composed by *Loxoconcha malcomsoni* Horne and Robinson, 1985, *Tuberoloxoconcha* cf. *atlantica* Horne, 1989, *Terrestricythere* cf. *elisabethae* Horne, Smith, Whittaker and Murray, 2004, *Xestoleberis labiata* Brady and Robertson, 1874 and many other more marine species such as *Basslerites teres* (Brady, 1869) and *Leptocythere fabaeformis* (G. W. Müller, 1894) (*e.g.* LOUREIRO *et al.* 2009; CABRAL & LOUREIRO 2013).

The knowledge of the influence of environmental parameters in the ecology of modern foraminifera, ostracoda, and many other groups, is a fundamental requirement to develop reliable palaeoenvironmental reconstructions.

6. CONCLUSIONS

The different climatic features recognized from the N to the S of Portugal, introduces significant differences in the hydrological balance of W coast estuaries. The contribution of freshwater drainage to the estuarine waters is reflected in the tidal marshes interstitial water parameters. The seasonal cycle is also well marked either the NW, under warm summers and no dry season (Cfb climate type) or the SW, under hot and dry summers (Csa climate type).

The tidal flooding waters that covers the salt marshes are not homogeneous. In the Minho low estuary the salinity may range between limnetic to euhaline conditions in the same tidal cycle. However, in the interstitial water mesohaline conditions prevail, under both the spring and autumn seasons, showing that tidal marsh may offer a relatively stable environment when compared with the extreme daily salinity variation of the lower estuary.

In contrast the Tejo and Mira salt marsh flooding waters tend to keep euhaline characteristics, although they can be slightly polyhaline in the rainy season. Such conditions and the control of evaporation/ precipitation balance are responsible by an enhanced euhaline to hypersaline interstitial marsh waters.

These environmental contrast are clearly reflected in the composition and distribution of tidal marsh assemblages. Namely foraminifera and ostracoda, that tend to be dominated by low salinity tolerant species in the low estuaries from NW and by high salinity tolerant species in the low estuaries of SW coast.

This data from Portuguese marshes represent

a reliable support for the interpretation of the geological record in a regional, as well as global context, of palaeoclimatic and palaeoenvironmental reconstructions of east North Atlantic estuaries.

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